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Morphometrical analysis of Pedhi River Basin Using GIS

Anand Rameshrao Dhote

In the present paper, an attempt has been made to study the quantitative geomorphology of Pedhi river basin it is sub-basin of Purna river of Maharashtra, India. Authors have evaluated the morphometric characteristics based on Survey of India toposheets at 1:50,000 scale, and LISS IIIimage. GIS and image processing techniques were adopted to identify the morphological features and to analysis their properties. Various morphometric parameters like linear and aerial aspects of the river basin were determined and computed. The results indicated that the drainage area of basin is 1460.30 Km2, perimeter 187.78Km, basin length 71.05 Km, stream order of the basin is up to Six order, stream length 2226.71Km., stream segments are 2027, Sinuosity Index is 1.41, Stream length ratio, mean Stream length, bifurcation Ratio, mean bifurcation ratio, is computed. Areal aspects are drainage density 1.53, stream frequency 1.39, texture ratio 10.79, elongation ratio 0.34, circulatory ratio 0.52, form factor 0.29, constant of channel maintenance 0.656 also computed. All the linear and areal parameters of the basin indicate gentle slope, with highly permeable bedrock. The relief ratio of the basin characterizes less resistant rocks. The outcomes thus generated equip us with significant knowledge and may provide an input that is essential in decision making for watershed planning and

y words: Quantitative geomorphology, Image processing techniques, Morphometric parameters. 1.0 Introduction:

'Morphometry may be defined as the measurement and mathematical analysis of the configuration of its earth's surface and the shape and dimension of it landforms'(J.I.Clarke1970). A major emphasis in geomorphology over the past several decades has been on the development of quantitative physiographic methods to describe the evolution and behavior of surface drainage networks (Horton, 1945). Morphometric analysis of a watershed provides a quantitative description of the drainage system, which is an important aspect of the characterization of watersheds (Strahler, 1964). Remote sensing and GIS techniques are now a day used for assessing various terrain and morphometric parameters of the drainage basins and watersheds, as they provide a flexible environment and a powerful tool for the manipulation and analysis of spatial information. In present study an attempt was made to analysis liner and areal aspects of Pedhi river basin with the help of

Location Map 88

Anand Rameshrao [

Pedhi is an important of Riddhapur 380 me and the west, the chie course passing by W westwards and northjourney outside the Ar Study Region:

The Pedhi w drainage basin of An latitude and 77°28' to 975 mm. The soils of possess clayey texture Objective: The main the basin.

Objective: The main the basin.

Materials And Metho Data and Methodolog In Present pa 55 H/5, and 9 and LISS

Attributes of various basins are computed through measurement No. 1.

ign measur	ement
Sr. No.	Pa
1	Strea
2	Strea
3	Stream
4	Mean
	(Lsm
5	Strea Ratio
6	Bifurc (Rb)

Mean Bifurcation Ratio (Rbm) Rbm = Average of bifurcation ratio of al order.	eference rahler (1957)
Areal Aspects Density (Dd) D = Where, D = Drainage density Lu = Total Stream Length of all orders A = Area of the basin (sq. km.) Frequency (Fs) Frequency (Fs) Texture Paris	rahler (1957)
Density (Dd) Density (Dd) Density (Dd) Density (Dd) Density (Dd) Lu = Total Stream Length of all orders A = Area of the basin (sq. km.) Frequency (Fs) Frequency (Fs) Frequency (Nu = Total Number of Streams of all order sA = Area of the basin (Sq. km) Texture Postion Density (Dd) Hotal Stream Frequency Nu = Total Number of Streams of all order sA = Areal of the basin (Sq. km)	
Density (Dd) Density (Dd) Density (Dd) Density (Dd) Density (Dd) Lu = Total Stream Length of all orders A = Area of the basin (sq. km.) Frequency (Fs) Frequency (Fs) Frequency (Nu = Total Number of Streams of all order sA = Areal of the basin (Sq. km)	
10 Stream Frequency (Fs) Frequency (Fs) Texture Post: Full Total Stream Length of all orders A = Area of the basin (sq. km.) Fs = Where,Fs = Stream Frequency Nu = Total Number of Streams of all order sA = Areal of the basin (Sq.km)	orton (1945)
Frequency (Fs) Stream Frequency (Fs) Fs = Where,Fs = Stream Frequency Nu = Total Number of Streams of all order sA = Areal of the basin (Sq. km.) Texture Posts	, ,
11 Texture Posice Areal of the basin (Sq.km)	
11 Texture Posice Areal of the basin (Sq.km)	orton (1945)
(Rt) Ratio Rt = Where Pt - T	
P Where, Rt = Texture Ratio Hot	rton (1945)
1 Otal Number of G.	11011 (1943)
ongation	
Ratio (Re) $Re = \frac{\sqrt{A}}{Where}, Re = Elongation Ratio$	umm (1956)
A = Area of the basin (Sq.km) $\pi = 'Pi' \text{ value } 3.14$	
l h = Posin Y	
(Rc) $Rc = Where Position$	er (1953)
$\mathcal{N} = 3.14 \text{A} = \text{Area of } 1$	ci (1953)
Ff = $\frac{A}{A}$ Where, Rf = Form Factor, A = Hort	ton (1945)
Area of the basin(Sq.km) $Lb_2 = Square of basin Length.$	

The various morphometric parameters such as liner and areal of the Pedhi river basin area were determined and are summarized in Tables No.1 and 2. The linear aspects of drainage network

The linear aspects of drainage network such as stream order (Nu), stream number (Nu), stream length (Lu), Mean stream length (km) (Lsm), Stream length ratio (RL), Bifurcation ratio (Rb), Mean bifurcation ratio (Rbm) result have been presented in the Table No.1 Table No. 1 Liner aspect of Pedhi river basin

River Basin	Stream Oreder	Stream	Stream Stream	Log Nu	17.
,	Oreder	number	length	Log Ivu	Log Lu
	1	(Nu)	In km.(Lu)		
D #1.	2	1522	1173.72	3.070	3.070
Pedhi	3	364	508.24 .	2.706	2.706
River	4	100	279.63	2.447	2.447
Basin	5	34	159.83	2.204	2.204
	6	6	32.91	1.517	1.517
		1	72.39	1.860	1.860

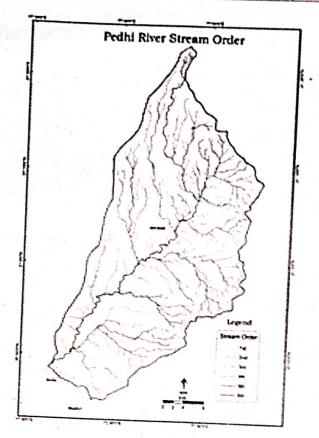
Stream Order (Nu):

In the drainage bas study, the channel segment ordering system. According order 1. Where two order 2 join, a segment of ord discharge of water and sedin is a 6th stream order of the ba Stream Number (Nu):

The law of stream no and stream numbers. Inverse increase number of stream is (1522), in 2ed it is 364, in 3rc decrease.

Anand Rameshrao Dhote

1stordar/	Bifur	cation Rat	io (R _b)	And the second second	Mean
2 ^{ed} order	3 ^{ed} order	3 ^{ed} order/ 4 th order	4 th order/ 5 th oSrder	5 th order/ 6 th order	Mean bifurcation ratio (R _{bm)}
2.309	1.818	1.750	4.856	0.455	2.240



Stream Order (Nu):

In the drainage basin analysis the first step is to determine the stream orders. In the present study, the channel segment of the drainage basin has been ranked according to Strahler's stream ordering system. According to Strahler (1964), the smallest fingertip tributaries are designated as order 1. Where two first order channels join, a channel segment of order 2 is formed; where two of order 2 join, a segment of order 3 is formed; and so forth. The trunk stream through which all discharge of water and sediment passes is therefore the stream segment of highest order. The study area

Stream Number (Nu):

The law of stream numbers relates to the definite relationship between the orders of the basins and stream numbers. Inverse correlation in stream order and stream number, as the order of stream is increase number of stream is decrease. In study area highest numbers of streams are in the 1st order (1522), in 2ed it is 364, in 3rd it is 100, in 4th it 34 and in 5th it is 36 6th is one decrease as the order is

1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 -
er (1957)
ton (1945)
on (1945)

on (1945)

mm (1956)

on (1945)

dhi river basin area

ream number (Nu), furcation ratio (Rb),

Log Lu
(C)
3.070
2.706
2.447
2.204
1.517
1.860

Stream Length (Lu):

Stream length is one of the most significant hydrological features of the basin as it reveals surface runoff characteristics streams of relatively smaller lengths are characteristics of areas with larger slopes and finer textures. Longer lengths of streams are generally indicative of flatter gradients, Generally, the total length of stream segments is maximum in first order streams and decreases as the stream order increases. The number of streams of various orders in the basin is counted and their lengths from mouth to drainage divide are measured with the help of GIS Software. Plot of the logarithm of stream length versus stream order (Figure 4b) showed the linear pattern which indicates the homogenous rock material subjected to weathering erosion Characteristics of the basin. Deviation from its general behavior indicates that the terrain is characterized by variation in lithology and topography. In study area number of stream is decrease as the order of the stream is increase.

The term bifurcation ratio (Rb) is used to express the ratio of the number of streams of an given order to the number of streams in next higher order (Schumn, 1956). Bifurcation ratios characteristically range between 3.0 and 5.0 for basins in which the geologic structures do not distort the drainage pattern (Strahler, 1964). Strahler (1957) demonstrated that bifurcation ratio shows a small range of variation for different regions or for different environment dominates. The mean bifurcation ratio value is 2.240 for the study area (Table 2) v hich indicates that the geological structures are less Areal Aspects of the Drainage Basin

Area of a basin (A) and perimeter (P) are the important parameters in quantitative morphology. The area of the basin is defined as the total area projected upon a horizontal plane contributing to cumulate of all order of basins. Perimeter is the length of the boundary of the basin which can be drawn from topographical maps. The aerial aspects of the drainage basin such as drainage density (D), stream frequency (Fs), texture ratio (T), elongation ratio (Re), circularity ratio (Rc) and form factor ratio (Rf) were calculated and results have been given in Table 3.

Horton (1932) introduced that the drainage density (D) is an important indicator of the linear scale of landform elements in stream-eroded topography. Drainage densities can range from less than 5 Km/Km2 when slopes are gentle, rainfall low, and bedrock permeable (e.g. sandstones), to much larger values of more than 500 Km/Km2 in upland areas where rocks are impermeable, slopes are steep, and rainfall totals are high (Huggett, 2003). The drainage density (D) of the study area is 1.53 Km/Km2. In the present study, the density falls less than 5 Km/Km2, which indicates that the area has a

Table 3: Aerial aspects of the study area

0	A Aerial aspects of the study area	,	
Sr. No.	Morphometric Parameters	D . 1]
2	Drainage density (D) (km/km2) Stream frequency (Fs)	1.53	
3	Texture ratio (Rt)	1.39	
4	Elongation Ratio (Re)	10.79	
5	Circulatory ratio (Rc)	0.34.	
6	Form factor (Rf)	0.52	
icy (Es).	- orm factor (KI)	0.29	

Stream Frequency (Fs):

Stream frequency or channel frequency (Fs) is the total number of stream segments of all orders per unit area (Horton, 1932). It exhibits positive correlation with drainage density in the watershed indicating an increase in stream population with respect to increase in drainage density. The

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Texture Ratio

Texture on the underlyin texture ratio of tl Drainage texture i) Coarse drainas ii) Intermediate iii) Fine drainage iv) Ultra fine drai

Elongation Rati Elongat an idea about the when the basin sh elongation, reachi value of 0.34 whic Circularity Ratio

Miller (19 area of circle havin The ratio is a l to a square, and antin Form Factor Ratio

The form t narrower basins hav and high form factor form factor for the ca gentle to moderately: factors greater than ur elongated basin. Conclusion:

Morphomatri through an early manu type drainage pattern. slope gradient. Lower underlined by uniform density of the Pedhi riv basin has a gentle slop texture, elongation ratio For the analysis fmon types of analysis frery Reference:

Chougale S. Maharashtra L 2017, Pg. 635-1 Horton, R.E.(Hydrophysical 370. In Olav Sla

Joji V. S., Naii Analysis of Par. GIS". Journal C

Stream Length (Lu):

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Areal Aspects of the Drainage Basin

Area of a basin (A) and perimeter (P) ard the important parameters in quantitative morphology. The area of the basin is defined as the/total area projected upon a horizontal plane contributing to cumulate of all order of basins. Perimeter is the length of the boundary of the basin which can be drawn from topographical maps. The aerial aspects of the drainage basin such as drainage density (D), stream frequency (Fs), texture/ratio (T), elongation ratio (Re), circularity ratio (Rc) and form factor ratio (Rf) were calculated and results have been given in Table 3. Drainage Density (D)

Horton (1932) introduced that the drainage density (D) is an important indicator of the linear scale of landform elements in stream-eroded topography. Drainage densities can range from less than 5 Km/Km2 when slopes are gentle, rainfall low, and bedrock permeable (e.g. sandstones), to much larger values of more than 500 Km/Km2 in upland areas where rocks are impermeable, slopes are steep, and rainfall totals are high (Huggett, 2003). The drainage density (D) of the study area is 1.53 Km/Km2. In the present study, the density falls less than 5 Km/Km2, which indicates that the area has a

gentle slope, moderate rainfall and permeable bedrock.

Table 3: Aerial aspects of the study area

	table Streeting appears	
Sr. No.	Morphometric Parameters	Result
1	Drainage density (D) (km/km2)	1.53
2	Stream frequency (Fs)	1.39
3	Texture ratio (Rt)	10.79
4	Elongation Ratio (Re)	0.34
5	Circulatory ratio (Rc)	0.52
6	Form factor (Rf)	0.29

Stream Frequency (Fs):

Stream frequency or channel frequency (Fs) is the total number of stream segments of all orders per unit area (Horton, 1932). It exhibits positive correlation with drainage density in the watershed indicating an increase in stream population with respect to increase in drainage density. The stream frequency value of the basin is 1.39

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Texture Ra

on the under texture ratio Drainage te: i) Coarse dri ii) Intermed iii) Fine drai iv) Ultra fin-Elongation

an idea abor when the ba elongation, value of 0.3 Circularity Mi

area of circl The ratio is a square, an Form Facts

Th narrower ba and high for form factor gentle to me factors grea elongated b Conclusio: Me

through an type draina slope gradi underlined density of t basin has a texture, elo For the ana types of anc Reference:

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Texture Ratio (T):

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Texture ratio (T) is an important factor in the drainage morphometric analysis which depends on the underlying lithology, infiltration capacity and relief aspect of the terrain. In the present study the texture ratio of the basin is 10.79 which falls within the fine drainage texture.

Drainage texture can be classified into four categories viz:

i) Coarse drainage texture: < 4

ii) Intermediate drainage texture: 4 - 10

iii) Fine drainage texture: 10-15

iv) Ultra fine drainage texture: > 15

Elongation Ratio(Re):

Elongation ratio is a very significant index in the analysis of basin shape which helps to give an idea about the hydrological character of a drainage basin. The elongation ratio varies from 1.275 when the basin shape is a circle, to 1.128 when it is a square, and decreases in proportion to increasing elongation, reaching a minimum of approximately 0.200. The results of the present study indicated a value of 0.34 which indicate elongation of the basin.

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Circularity Ratio(Rc):

Miller (1953) defined a dimensionless circularity ratio (Rc) as the ratio of basin area to the area of circle having the same perimeter as the basin. Circulatory ratio of the present study area is 0.52. The ratio is equal to unity when the basin shape is a perfect circle, decreasing to 0.785 when the basin is a square, and continues to decrease to the extent to which the basin becomes elongated, Miller (1953).

Form Factor Ratio (Rf)

The form factor will be comparatively higher, if the basin is wider. Consequently, much narrower basins have low form factor values. The low form factor is indicated in the elongated basin and high form factor is indicated in the wider basin (Gregory & Walling, 1985). The calculated value of form factor for the catchment is 0.29 Sparse to dense vegetation along with some current jhumming, gentle to moderately steep slope and low to high relief characterize the low value of form factor. Form factors greater than unity are considered as anomaly in the basin shape. The value indicates narrow and elongated basin.

Conclusion:

Morphomatric analysis of Pedhi river basin shows 1st to 6th order of stream, and it is passing through an early mature stage to old stage of the fluvial geomorphic cycle. The basin shows dendritic type drainage pattern. The larger number of first order streams points to uniform lithology and noble slope gradient. Lower value of bifurcation ratio indicates that the drainage of Pedhi river basin is underlined by uniform materials and the streams are usually branched systematically. The drainage density of the Pedhi river basin is 1.53 Km/Km2. It is less than 5 Km/Km2, which indicates that the basin has a gentle slope, moderate to low relief. Texture ratio of basin falls within the fine drainage texture. elongation ratio and circularity index and form factor shows that the basin is elongated shaped. For the analysis of morphomatric parameters topographical maps and LISS III image are used and this types of analysis is very useful for watershed planning and management.

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> *Anand Rameshrao Dhote Department of Geography, Yuvashakti Arts and Science College Amravati

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1. Introducti

The natural surro population w distribution. · understandir regional plar every countr population g has toughen decentraliza crucial facts developmen favourable f of land is his conditions s It is estimate this popular projections will not only on the avail population: is inversely growth (Co factor to au; growth alor 2015) and c ahead to ha